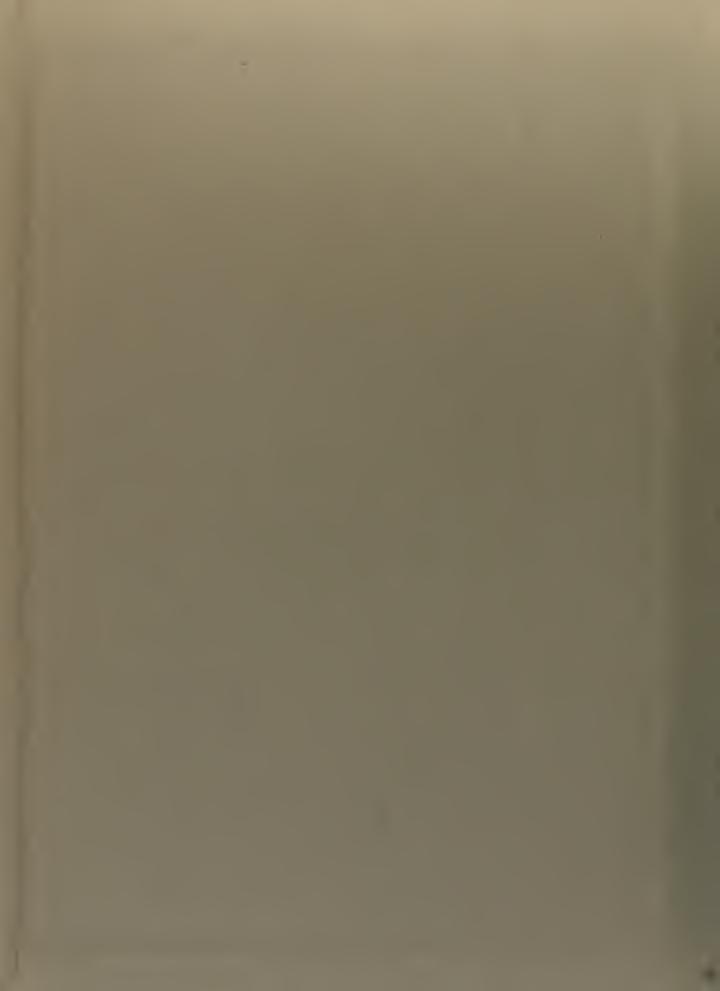
A STUDY OF THE INTERRELATION OF POROSITY, PERMEABILITY, AND GRAIN SIZE IN A CONSOLIDATED POROUS MEDIUM

Strart Allan MacCaffray











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Stuart Allen MacCaffray B.S. in Ch.B., Northeastern University, 1941

Submitted to the Graduate School of the University

of Pittsburgh in partial fulfillment of the

requirements for the degree of

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POSEWORD

the gas and all industry of today is still plagued with unresolved and unexplained problems and phenomena, ospecially in the field of reservoir engineering. One of these is the macroscopic flow of fluids in reservoir rock formations where gas and oil are found and produced. This thesis is a study of but a small facet of the problem and has been conducted with the idea that any work done in this area will help contribute to an increased knowledge of the subject.

The author wishes to acknowledge the sincere and helpful suggestions, advice, and encouragement of Professor Holbrook G. Botset, Head of the Petroleum Engineering Department, University of Pittsburgh, in the projection, progression, and successful completion of this study.

Acknowledgment is also made to those authors, investigators, and researchers whose works were utilized for background and basis of the study without which such an undertaking would have been impossible.

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I. INTRODUCTION

and oil, far beneath the surface of the earth in areas of rock formation called reservoirs. These reservoirs are not, as is usually presumed by the uninformed, large voids filled with gas and oil which, when tapped, may be easily produced. On the contrary, they are formations of porous and permeable rock with small, macroscopic, open spaces between the grains of the matrix material in which the gas and oil is stored; a factor which makes the production of such stored gas and oil difficult and, in some cases, impossible. It is for this reason that the complete study of all conditions and probabilities of fluid flow in porous media becomes an important facet in the production of gas and oil.

evaluate the inter-relationships of some of the major physical properties of porous reservoir rock which affect the flow of fluids through it. Although these inter-relationships have been expressed by several different formulations, mentioned later, all designed to elucidate fluid flow, it was thought that a more useful and informative modification could be developed and applied to the problem of correlating permeability, porosity and grain size of porous media to fluid flow through such media. Such a study, designed to reveal the mechanics of macroscopic fluid flow in porous media, should be useful in evaluating some of the many factors which must ultimately be known if a satisfactory solution of the problem is ever to be obtained. Any additional knowledge of these factors should contribute to a more complete understanding of the complexities of gas and oil production from the reservoir.

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engineering, a very important physical property of reservoir rock is permeability, which can be properly defined as a measure of the fluid-transmitting capacity of a porous material, or as the ability of a fluid to flow within the interconnected pore spaces of a porous material. A second physical property, also equally important, is porosity. In oil and gas reservoirs, porosity represents the percentage of total rock volume which is available for occupancy by either liquids or gasses or both. In connection with the term porosity, care should be taken to distinguish between absolute and effective porosity. Absolute porosity is the percentage of void space with respect to total volume. Effective porosity, on the other hand, is the percentage of interconnected void space in the total volume. A third physical property of reservoir rock is actual grain size and grain size distribution of the matrix material which forms the reservoir rock.

that permeability, effective porceity, and grain size are interrelated, one depending upon the other to varying degrees. In porcus material in which the porce are inter-connected, there exists no definite relation between permeability and effective porceity. These, however, are directly influenced by grain size, in that grain size or particle size and its attendant factors of uniformity of grain size, shape of the grain, manner of grain packing, and the amount of cementing material between the grains (degree of lithification) determine the size and shape of the pore openings, the extent of interconnections of the pore spaces, and the ratio of their volume to the total volume of the rock.

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Investigations^{1,2,3,6,10} have indicated that actual grain size and effective porosity are excellent criteria for judging permeabilities of various porous media. In much of the literature, variations of permeability and porosity within individual reservoirs have been found to be inter-related to a degree and have been shown by Fettke^{1,2} in the case of the Bradford Third sand of the Bradford field of Fernsylvania to have these relationships

and

where K is the permeability and \$\beta\$ is the perceity. The first formula was determined from earlier experiments and the second from later experiments. However, these relationships are average and do not apply in some cases or throughout the complete possible range of permeability and perceity. Here again grain size, with its attendant factors, directly affects these relationships.

In making determinations of the values of permeability, effective perosity and grain size diameter, the first two are now very easily and effectively carried out by means of simplified, standardized methods which give fairly accurate results. In the case of average grain size diameters, this determination becomes complicated by the particle itself, in that the particle is usually irregular in size and shape, and no known methods are available for defining such a particle in geometric terms. This problem becomes more complicated and formidable in the case of an aggregate of irregular particles, but has been successfully solved by the use of statistical methods for evaluating surface area and volume of the aggregates.

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Such mothers are based on the measurement of the so-called average grain diameter by relating the irregular particle to an equivalent geometric figure such as a circle or a sphere from which a diameter can be calculated. These methods determine surface area and volume from the number and weight of the particles, using as a basis the theory of disensions which requires that the surface area be proportional to the square, and the volume be proportional to the cube of the dimensions used.

For the determination of particle size to have physical significance, that is, to relate particle size to definite physical properties such as surface or volume, the arithmetic or geometric mean and the median dismeter determinations are supermeded by statistical diameters. Several formulas have been developed and used to designate these statistical diameters in terms of various physical properties such as mean diameter (length), mean volume diameter, mean surface diameter, mean volume-surface diameter, and weight mean diameter. The average grain diameters used in this study are mean volume diameters, based on the average volume and defined to be that diameter whose volume divided into the total volume would give the number of particles. If d is the average particle diameter and V is the total volume, a relationship exists as

where n is the number of particles of a given size. If d_v^3 represents the volume of an average particle, and $\geq n$ is the total number of particles then

$$n = \le nd^3/d_{\psi}^3$$

 $d_{\psi} = (\le nd^3/\ge n)^{1/3}$

and

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II. HISTORICAL BACKUNOUND

Investigation of the flow of fluids through porous material at low pressures was initiated by Parcy³ in 1856. These studies consisted of a series of experiments on the flow of water through filter beds from which Parcy obtained an empirical formula showing the rate of flow to be very nearly proportional to the pressure drop per unit length of porous medium. In its elementary form, this formula,

indicates that the fluid permeability, K, of a porous material proportionately relates the velocity, v, of flow of a fluid of μ viscosity to the pressure differential, Δ P, causing this flow over a length of Δ L. Since Δ P is measured between the outflow and inflow ends, it is negative as indicated in the formula above.

The proportionality constant K is considered to be a specific property of the porous material, empirically independent of the dimensions of the material, the pressure differential exerted on the fluid flowing, and the viscosity of the fluid. For this reason, K may be expressed in terms of other measurable physical properties of the porous material, such as porosity and grain size. Since the formulation of Parcy's equation and from it, his law, many investigators and researchers have attempted to show relationships of porosity, permosbility, and grain size.

In 1880, Seelheim¹⁷ first introduced grain size into the relationship by showing that the rate of flow was proportional to the square of the average grain dismeter. This would indicate a decrease in permeability as sand grains decrease in size. and in the control of the control of

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The next investigator to utilize grain size in flow relationship formulas was Hazen who, in 1592 and 1893, proposed the formula based on studies of sand filters,

$$q = E_{\rm H}(d_{\rm H})^2(h/L)$$

where $E_{\rm R}$ is a permeability constant, h/L is the pressure drop through the filter per unit-depth, and $d_{\rm R}$ is the "effective" dismeter of sand particles having uniformity coefficients of less than 5. "Effective" grain size is defined as the opening which will just pass ten per cent of the particles (by weight) and uniformity coefficient is defined as the ratio of the size opening which will pass 60 per cent of a sample being screened, to the size which will just pass 10 per cent. When the uniformity coefficient is low, as 5 or less mentioned above, the particles are more or less uniform in size, but as the uniformity coefficient increases to higher values, the particle sizes become widely distributed.

Water per day per square foot of filter area, d_H in millimeters, the loss of head, h, in feet of water, and L in feet, then the values of E_H range from 1300 to above 4000 depending on the cleanliness of the sand, with the usual limits of E_H for ordinary sands ranging between 2300 and 3300. Hazen's equation has been widely accepted and utilized by sanitary engineers and is of considerable historical interest since it represents the first attempt to recognize the importance of particle size and a method for its satisfactory representation. Mavis and Wilsey⁵ retested Hazen's formula with the following results:

$$q = 2300 d_{\rm H}^2 \left(\frac{T + 10}{50}\right) (h/L)$$

 $q = 780 (d_{\rm hh})^2 (d/h0)^6 \left(\frac{T + 10}{50}\right) (h/L)$

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where T is the temperature of the water in degrees Fahrenheit, dal is the

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screen size in millimeters through which just 34 per cent of the particles (by weight) pass, and \emptyset is the porosity introduced into the second formula.

Slichter, in 1899, published the accounts of his theoretical analysis of fluid flow through an ideal homogeneous packing composed of apheres of uniform size in which he was the first to introduce the effect of packing as a factor influencing permeability, which gave rise to expressing the average pore area in terms of the diameter of the sphere.

Slichter derived a modified Poiseuille relationship in his equation

where q is expressed in cc of water per second, A is the cross-sectional area of the packing in square centimeters, h is the pressure differential in centimeters of water, d is the diameter of the spheres in centimeters, K_S is a packing constant dependent on porosity, a is the viscosity of the fluid flowing in poises, and L is the thickness along the direction of flow in centimeters. In terms of Darcy's permeability, Slichter's formula would read

$$K = 10.2d^2/K_{\rm s}$$

For various packings, tabulations have been made of the different values of the packing constant K, shown as a function of porosity for porous material packed with spheres of equal dismeter.

Havis and Wilsey⁵ developed a formula $1/\pi_a = 0.05(6/40)^{3.3}$

for use in the determination of the values of K_s where \$\psi\$ is the perosity. King recognised the importance of Slichter's formula by using it to calculate the average diameter of irregular particles. It is important to note here that the average diameter of irregular particles determined in this manner is the diameter of the particles such that if all of them were

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of this diameter, the packing would have the same permeability as it actually has at a given temperature and porosity.

For the ultimate in permeability, porosity, and grain size reletionships, it remained for Kosery, 8 in 1927, to develop and publish his important empirical formula

where C is a form factor dependent on the shape of the pores, \emptyset is perceity, g is the acceleration constant, \triangle P is the pressure differential in grams per square centimeter, u is the viscosity of the fluid flowing in poises, S is the specific surface area in square centimeters per cubic centimeter, L is the thickness of the porous medium in centimeters, L₀ is the longer path of flow in centimeters through the thickness L, and L/L_0 is a reduction factor for the relationship \triangle P/L. Since S a S₀(1 - \emptyset) the above equation can be rewritten in terms of S₀, which is the specific surface area per unit volume of matrix material, cm⁻¹, as

Unaware that Kozeny had developed the aforementioned relationship, fair and Natch, 9 in 1933, developed an almost identical formula by following a line of reasoning similar to that pursued by Kozeny. This equation is

It can be seen that the factors (C.L/La) in the Kozeny equation have been replaced by the constant 5, believed by Fair and Hatch to apply to normal, rather compact, unconsolidated percus material.

Carmen 10,11 in 1937, presented an improved derivation of the Moveny and Pair and Hatch equations. Using as a basis the general equation for flow through pipes

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where $\mathbf{v}_{\mathbf{c}}$ is the average velocity, g is the acceleration constant, m is the hydralic radius, defined as the volume of fluid in the pipe per surface area in contact with the fluid, Δ P is the pressure differential, $k_{\mathbf{c}}$ is a shape factor and is the reciprocal of Kozeny's form factor C, u is the viscosity of the fluid flowing, and $\mathbf{L}_{\mathbf{c}}$ is the length of flow path, Carmen set about to substitute or modify the equation factors so that the equation would represent flow of fluid through perous media. For v he substituted \mathbf{v}/\mathbf{v} , later to modify it by the relationship $\mathbf{L}_{\mathbf{c}}/\mathbf{L}$ for flow over the flow path, $\mathbf{L}_{\mathbf{c}}$, greater than the k dimension of the medium. Then he substituted for m the relationship of $\mathbf{v}/\mathbf{s}_{\mathbf{c}}(1-\mathbf{v})$, and applying the reducing factor $\mathbf{L}/\mathbf{L}_{\mathbf{c}}$ to the expression Δ P/L, he arrived at the following rearrangement:

In a series of experiments using a very wide range of particle shapes in unconsolidated media, Carsan found that the parameter $(C \cdot L/L_e)^{-1}$ of Hozeny's equation and the parameter $k_0(L_e/L)^2$ of his equation both averaged about 5. In the modified Hozeny-Carsan equation these parameters are expressed as k, with an assigned value of 5, shown by

Wyllie, 12 in his paper on the historical development of the Kozeny-Carman equation, discusses the more recent determinations by investigators of values greater than 5 for the Kozeny-Carman constant k when the formula has been used in connection with measurements of flow through consolidated modia. Since the k_0 shape factor of the Kozeny-Carman equation has been held by most investigators to be constant, the higher values of k have been attributed to the $(l_0/L)^2$ factor which has and

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Bruce¹³ have called the tortuosity factor, T. Several determinations were made by Mose and Bruce for the value of k for various consolidated media of low permeability where values were obtained as high as 100.

Others concur with Mose and Bruce, all indicating that the constant for consolidated media varies to higher values and that these higher values are attributable to the tortuosity function.

There has been some disagreement with the concept of increasing values of the Koseny-Carman constant in connection with consolidated media. H. E. Roselh believes that errors arising from the use of 5 as a value for the constant will be far loss than those arising from uncertainties in other variables, especially the porosity function, expressed as $\beta^3/(1-\beta)^2$ in the Koreny-Carman equation. He has shown, by plotting relative resistance versus per cent perceity, that at perceities lower than 40 per cent there are increasingly different values of relative resistance to flow for various equations which all contain a porosity function of some mort. Valla Valle 15 has shown a similar plot of relative permeability versus per cent perceity for several different investigators' formulas, all containing a porosity function. This also indicated that at porosities lower than hO per cent the perceity function does not remain constant. It has not been possible to test the porosity function $f^3/(1-g)^2$ in consolidated media because it is not readily possible to vary peresity and keep the specific surface area constant.

A recent modification has been made in the Kozeny equation by Wyllie and Spangler 16 who have combined the Kozeny equation with properties of the capillary pressure desaturation curve to obtain the formula

$$K = (\gamma^2/2.5r^2\phi) \int_0^1 ds_y/r_0^2$$

where K is the parmeability coefficient, y is the interfacial tension, F

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is the formation factor, \emptyset is the porosity fraction, S_w is the fractional vetting phase saturation, and P_c is the capillary pressure. This equation introduces new physical factors which are more easily measurable in laboratory determinations and consequently should further the continuing search for more accurate and useful relationship formulations for fluid flow through porous media.

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III. PHOCEDINE

and calculated values for permeability, together with accompanying values for porosity and average grain size, were selected from Fettke's 1,2 studies of the Bradford Oil Field, Bradford, Pennsylvania. These data had been derived from test determinations made on representative core samples obtained from various sections of the field at random depths in the Bradford Third sand. The Bradford Third sand is an excellent representative of a consolidated porous medium whose general characteristics are low permeability and porosity and quite fine grain size with irregular shape and size distribution and a moderate to high amount of lithification. Since practically all oil and gas reservoirs are located in consolidated rock formations, a representative sample of this type of porous medium was chosen for application to this study.

As a first approach to the problem, an investigation was made of permeability, peroxity and grain size relationships when applied to an equation modified by the author from one developed by Traxler and Baum from the Poiseuille equation of flow through pipes. This equation was expressed by Traxler and Baum as:

$$D_{\rm e} = (32q_{\rm A}L/Ag\Delta P)^{\frac{1}{2}} \tag{1}$$

where $D_{\mathbf{c}}$ is the effective pore diameter in centimeters, q is the rate of flow in cubic centimeters per second, μ is the viscosity of the fluid flowing in poises, L is the length of flow path in centimeters, A is the cross-sectional area in square centimeters, β is the perceity, and Δ P is the pressure differential in dynes per square centimeter. It can be readily seen that the function $\alpha + 1/A \Delta$ P of this equation will be equivalent

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to the permeability coefficient K, in darcies, when the units are changed to those of the Darcy equation. Thus, K = (ql00 μ L/A9.8692 X 10⁻⁷ Δ F) or 9.8692 X 10⁻⁹ E = q μ L/A Δ P. Substituting in equation (1) above gives:

$$D_{a} = 5.619 \times 10^{-10} (K/8)^{\frac{1}{2}}$$
 (2)

where K is the permeability coefficient in darwies.

Table I, Appendix I, has been presented to show calculations which resulted from the application of the permeability and poresity data selected for this study to equation (2) above, solving for $D_{\rm G}$, the effective average pore diameter. Sutting 19 has stated that the effective average pore diameter is roughly one-fifth of the average grain diameter. Figure 1, Appendix II, shows a plot made from the data in Table I of d, the average grain diameter, versus $D_{\rm G}$, the effective average pore diameter, resulting in a straight line whose slope, which represents the ratio of $d/D_{\rm G}$, is approximately 25. The value of this ratio is considerably greater than the one of 5, $(1/5d \pm D_{\rm G})$, proposed by Sutting. It should be explained here that only the results of the calculations using Fettke's 1934 data were used in plotting Figure 1. The results of the calculations using Fettke's 1933 data produced no conclusive plot.

As a second approach to the problem, an investigation was made of permeability, porosity and grain size relationships when applied to a formulation derived by equating the Kozeny-Carman and the Darry equations. This second study was made on the basis of the specific surface area, S_0 , of percus medium per unit volume of matrix material. By taking the basic Kozeny-Carman equation solved for S_0^2 one obtains:

$$s_0^2 = \frac{gg^3 A \Delta P}{5(1-g)^2 \text{ Qub}}$$
 (3)

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It can be seen that in equation (3) on the preceding page the function $(A\Delta P/quL)^{-1}$ is equivalent to the K of Parcy's equation when the units of μ and ΔP are changed from poises to centipoises and from grans per square centimeter to atmospheres respectively. Thus, K = $qLOO_{cl.}/(A9.678 \times 10^{-4}\Delta P)$ or 9.678×10^{-6} K = $(A\Delta P/quL)^{-1}$. Substituting in equation (3) gives:

$$s_0^2 = \frac{s_0^3}{5(1-s)^29.673 \times 10^{-5}R} \tag{4}$$

Since the acceleration constant, g, is 930 centimeters per second equared, equation (h) above can now be resolved as

$$S_0 = 4501 \cdot \frac{6}{1-6} \cdot (4/8)^{\frac{1}{2}}$$
 (5)

which is the equation used in this second approach to the problem under study.

Table II, Appendix I, has been prepared to show calculations resulting from the application to equation (5) above of the permeability and porosity data selected for this study. It has been shown by Herian that $S_0 = Z_g/Z_V$ for irregular particles where Z_g/Z_V is the ratio of the surface to volume factors and d is the average grain diameter. In this ratio, $Z_g = S/d_g^2$, where S is the surface area and d_g is the statistical particle diameter based on surface area, and $Z_V = V/d_V^3$, where V is the volume and d_V is the statistical particle diameter based on volume. The ratio Z_g/Z_V has been indicated by Fair and Hatch to be a useful measure of particle shape and has been given a value of 6.0 for spheres, 6.1 for rounded particles, 5.4 for worn particles, 7.0 for sharp particles, and 7.7 for angular particles. Figure 2, Appendix II, shows a plot made of

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the values in Table II of S_0 versus 1/d which roughly form straight lines with a slope of 143 for line 1 using Fettke's 1934 data and a slope of 59 for line 2 using Fettke's 1938 data. The slope normally should represent the values of the ratio Z_1/Z_2 which lie somewhere between 6.0 and 7.7.

As a third approach to the problem, an investigation was made of permeability, perceity and grain size relationships when applied to a formulation derived by equating the Koseny-Carman and the Barcy equations, as previously mentioned, except that the formulation was modified for the use of average grain size in lieu of specific surface area. Since S_0 is proportional to 1/d, S_0^2 can be replaced by $1/d^2$ in equation (h) above. There is also a new constant, k', introduced into the equation which is equivalent to the function $1/(2g/2_{\psi})^2 S$ where S is the Komeny-Carman constant and $2g/2_{\psi}$ is the ratio of the surface to volume factors. The resulting equation is expressed as:

$$\frac{1}{d^2} = \frac{k^4 z g^3}{(1 - g)^2 9.678 \times 10^{-6} g}$$
 (6)

By substituting 980 centimeters per second squared for g in equation (6) and rearranging to solve for k*, the equation can now be expressed as:

$$k' = \frac{K(1-\phi)^2}{d^2\phi^3 1.013 \times 10^3} \tag{7}$$

Table III, Appendix I, has been prepared from calculations derived by applying the data used in this study to the factors of equation (7); namely, K/d^2 and 1.013 \times $10^8 \beta^3/(1-\beta)^2$. The values of these factors were plotted as K/d^2 versus 1.013 \times $10^8 \beta^3/(1-\beta)^2$, shown by Figure 3, Appendix II, to be a straight line with a slope of 0.0001052 in the case of line 1 using Fettke's 1934 data and a curve with a variable slope of

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unknown value in the case of line 2 using Fettke's 1938 data. Line 2 could not be expressed as a straight line by repletting on semi-log or log-log graph paper. The slope of line 1 of Figure 3 is equal to the constant k' of equation (7) on the preceding page.

As previously stated, k' is equivalent to the function $1/(2_g/2_v)^25$. Using the value of the slope determined for line 1 of Figure 3 for k', the value of $2_g/2_v$ as obtained from the relation $2_g/2_v = (1/k!5)^{\frac{1}{2}}$ is 13.5. Normally, values for $2_g/2_v$ lie between 6.0 and 7.7 as previously montioned.

In the two preceding investigations, the perceity function, $\beta^3/(1-\beta)^2$, although it varies with the value of perceity, is considered a constant factor in that it retains the same fixed relationship regardless of the source of data. For a fourth approach to the problem, an investigation was made of the permeability, perceity and grain size relationship when applied to a formulation developed to express the perceity function as a variable, dependent upon the source of data, with the permeability and grain size being constant factors regardless of the source of data. Thus, the perceity function would be a constant factor only for a particular reservoir formation or even a particular section of the reservoir formation.

To develop this formulation, the basic Kozeny-Carman equation was recast to include particle diameter, giving:

$$q = \frac{g k! A g^3 d^2 \Delta P}{\mu L (1 - g)^2}$$
 (8)

whose units have been previously defined. Solving for the porosity function gives:

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$$\frac{g^3}{(1-g)^2} = \frac{q \mu L}{A \Delta P g k! d^2}$$
 (9)

Since the function $q_{1L}/A\Delta P$, when expressed in terms of the Parcy equation, is equal to the permeability coefficient K of Parcy's equation, substitution of 9.673 X 10-6K for $q_{1L}/A\Delta P$ can be made in equation (9) above giving

$$\frac{d^3}{(1-d)^2} = \frac{9.678 \times 10^{-6} \text{K}}{8 \text{ K}^4 \text{ d}^2} \tag{10}$$

Assuming that the constant k* would be reflected in a recasting of the porosity function, a new expression of this function can be substituted for $k! \frac{d^3}{(1-d)^2}$ in equation (10) above giving

$$(\emptyset)^{n} = \mathbb{E} / 1.013 \times 10^{8} d^{2}$$
 (11)

or, by rearrangement,

$$E = 1.013 \times 10^8 \text{ d}^2 \text{ (g)}^{23}$$
 (12)

where n is a constant applicable only to a particular formation or section of a fernation, or a porous medium.

Table IV, Appendix I, has been prepared from calculations derived from the application of Fettke's 1934 data to the functions of equation (12) above. Figure 4, Appendix II, is a log-log plot of these functions, $K/d^2/1.013 \times 10^8$ versus \emptyset , which forms a straight line with a slope of S. Since

$$n = \frac{\log K/d^2/1.013 \times 10^8}{\log \beta}$$

then n = 0, the slope of the straight line.

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IV. COM ENT AND TISCUSSION OF BESULTS

Before discussing the results, it should be pointed out here that the values of porosity data used in this study represent absolute porosity. However, Fettke² has shown that in a majority of cases the difference between absolute porosity and a lower value for effective porosity in the Bradford Third sand is but a matter of 0.1 per cent average. There are a few isolated cases of a deviation slightly greater than 0.1 per cent up to about 1.0 per cent. For purposes of this study, the values of effective percently were assumed to be the same as those for absolute percently. This fact can explain some of the slight deviations of the points used to plot the various lines and curves shown by Figures 1, 2, 3 and 4. Other deviations of the points can best be explained by the fact that at the time the data were generated, the accuracy of the determinations of values for permeability, percently and grain size was not nearly as good as that of the present day, due to greatly improved and standardized methods for determining such values.

the average pore disaster in a rather fine-grained, consolidated medium decreases considerably from the average for unconsolidated media, stated by Butting 19 to be roughly one-fifth of the average grain dismeter. This can be best explained by the non-uniformity of particle size and shape, the fine grain size, and by the degree of lithification or communing found in consolidated media. It is considered by the author that the latter exerts the greatest effect upon average pore dismeter since comenting to any extent tends to restrict, reduce and plug the pore channels of the rock. To utilize this method of correlation of permembility, percently

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and grain size, it would be necessary to determine the relationship between average effective pore diameter and average grain diameter for each particular reservoir or section of the reservoir.

In the second approach to the problem, the results of the investigation indicate that the value of the Kozeny-Carman constant for consolidated media is such higher than the assigned value of 5. By way of example, presume the consolidated medium to be composed generally of sharp particles where the value of the ratio of surface to volume factors, Z_g/Z_v , is 7.0. It has been previously shown that $S_o = Z_g/Z_v$ d. Weing as a selection from Fettke's 1934 data, d = .0093 centimeter, h = .0034 Parcy, and Ø = .145, it is possible, by resolution of equation (4) of section III of this study, to solve for a value of the Roseny-Carman constant in lieu of the value of 5 normally used. 5, is resolved as 7/.0093 or 753. Solving for the unknown Kozeny-Curman constant using the above values in the equation gives a value of approximately 220. By using some of the other sets of values for perseability, porosity, and grain size from the 1934 data to solve for the Koweny-Carman constant in the same manner as described above gives values for the constant in the general vicinity of 220, with some variations. The results of similar application of Fettke's 1938 data were not as conclusive, with the values showing a much wider variation, 30 to 107, for the Kozeny-Carpan constant.

As in the second approach, the third approach to the problem also indicated that the Koseny-Carman constant for consolidated media is much higher than 5. It has been previously stated in Section III of this study that the slope of line 1 of Figure 3, Appendix II, was 0.0001052 and equivalent to the function $1/(2_8/2_V)^2$ k where k is the Kozeny-Carman constant. If, for example, it is again assumed that $2_9/2_V$ has a value of

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7.0 for this medium, it is possible to solve for a value of other than 5 for the Kozeny-Larman constant. In this case, the value we resolved as 194 for the constant. Since a small deviation in the drawing of line 1 of either Figure 2 or 3 results in a change in the value of the slope of the line, it can be presumed that this fact caused the difference in values for the foseny-Carman constant determined by the two investigations. Therefore, it can also be presumed that in the case of the 1934 data used the average value for the Kozeny-Carman constant lies according between 194 and 220. We similar success was achieved in using Fettler's 1938 data. It is believed that the 1933 data were not truly representative of the reservoir sand nor were they in themselves complete enough to permit drawing any conclusions.

perceity function as a variable rather than a constant factor in lieu of a variable for the Kozeny-Carman constant as shown by the second and third investigations. An application of the 1934 data to equation (12) of part III indicates that the n of the perceity function, \$\mathscr{g}^n\$, is approximately equal to 3, as determined from the slope of the line plotted in Figure 4, Appendix II, with higher and lower variations in some cases. Since the perceity function is an arbitrary selection by the author, it follows that this particular function, being an exponential, could have a variety of forms all of which would, when solved numerically, give substantially the same values. \$\mathscr{g}^n\$ was adopted because of its simplicity if for no other reason. Although this method shows that the perceity function can well be a variable factor, it is believed that the works of Some and Bruce and others have conclusively shown that the perceity function remains constant while the Kozeny-Garman constant becomes a variable factor in

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consolidated media. No attempt was made in the fourth investigation to make use of Fettke's 1938 data because of provious inconsequential results.

In all of the author's four investigations, it has been shown that the procedures, equations, and constants applie ble to the empirical formulations derived from the studies of unconsolidated porpus media do not hold true in the case of consolidated percus media. However, it also has been shown that for a given or particular consolidated porcus rock formation, the empirical formulations apply when the normal constants are changed to those values which fit the situation, as in the case of this study of the consolidated Bradford Third sand. The evidence does not conclusively prove whether the porosity function should be a variable or a constant factor in permesbility, perosity and grain size relationship formulations. Nuch more investigation and research must be carried out along these lines before a conclusion can be reached as to which theory is correct, although as mentioned in the preceding paragraph present evidence points to the poresity function being a constant factor. It is very possible that both the porosity function and the Kozeny-Carsan constant might be found to be variables in the case of consolidated porous andia.

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V. CONCLUSIONS

sive, complex and relatively unknown. For these reasons, the study of particle size in relation to permeability and porosity of perces media is, of itself, extensive, complex and relatively unknown, especially in the case of consolidated porous media. The author agrees with Muskat²¹ who stated, in effect, that the use of small particle statistics and measurements in determinations of permeability was so complex and unresolved that the presently utilized methods of determining permeability and porosity far outweigh any method using grain size or its attendant properties to determine same. However, the author feels that it is essentially important to intensively study the field of particle size in order to gain a thorough knowledge and understanding of macroscopic fluid flow in percus media in the light of grain size and its attendant factors of packing, uniformity of size and shape, and degree of lithification, all of which affect fluid flow in such media.

Through an intensive study of this nature, it is possible that some day a formulation will be developed which will truly hold under all conditions and in any circumstances whereby the physical properties of permosability, porosity, and grain size with its attendant factors, can be successfully correlated. The four methods of correlation employed in this study can be used in a restricted degree when applied to the consolidated porous media of particular reservoirs or of sections of the recervoirs only after considerable research to determine the proper constants applicable in each case. In the same vein, and of considerable interest, is the method developed by Ryder²² for making permeability

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developed formula, good only for one particular sand, which uses only determined weights of sloved sand particles which have been passed through a series of U. S. Standard sleves.

Another possible investigation or method, suggested to, but untried by the author, would be to plot the results of sleving data from a core sample (per cent by weight of sample through and on each screen size) versus the average screen size for that per cent by weight. From the resultant curve, select a "representative" per cent value and determine the corresponding average screen size or presumed average grain diameter. Using equation (2) of Section III of this study, determine which per cent by weight value and corresponding screen size or presumed average grain size can be used to satisfy the equation, and then determine whether or not this same per cent value can be used in connection with all core sample data from cores taken from the same formation or section of the formation.

A summation of the results of this study does not reflect the original purpose of the study. It was thought that an evaluation could be made of the interrelationships of the physical properties of porous rock, permeability, porosity and grain size, as well as a modification of existing formulations used to express these interrelationships of physical properties. By using data from highly consolidated porous media, it was hoped that if a modification were developed for effectively expressing the interrelationship of permeability, porosity, and grain size, it would hold under any circumstances and conditions for any type medium, either consolidated or unconsolidated. However, the complexities of grain size study, the learth of sufficient and accurate data, and the apparent non-conformity of application of modified formulations to consolidated porous

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media data made such endeavors impractical and impossible.

It is believed that the results of this study, although meager, can contribute some benefits in a better unterstanding of the problem. The ratio of pore volume to grain size diameter has been shown to decrease markedly with an increase in cementation and a corresponding decrease in grain size such as found in a consolidated perous media like the Bradford Third sand. The value of the Kozeny-Carsan constant has been shown to increase to very high values in consolidated porous media. This is probably due to the large increase in values of the tortuosity factor which becomes greater as the rock becomes more consolidated because of an increase in the effective length of flow path. Probably the most important contribution of this study is the realization of the importance of the Kozeny-Carman equation and its modifications when applied to problems of fluid flow through porous media. It helds great possibilities, when subject to the proper modification, in its application to studies of this type and gives promise of eventual solution of the many problems of fluid flow still facing the petroleum industry.

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APPENDIXI

Fettke's 193h Tata (N is Experimental)

(1)	(2)	(3)	(4)	(5) 1	(6)	(7)
Grain Size	Permeabili ty	Porosi ty	1./4	$(K/\emptyset)^{\frac{1}{2}}$	Constant	0c = C(R/Ø)
d (ca)	K (farcies)	Ø (per cent)			x 10-4	(en) x 10-5
.0092	.0005	10.8	.00460	.0578	5.619	3.82
.0092	.0005	12.5	.00400	.0633	0	3.56
.0093	.003ls	14.5	.02340	.1530	ti	8.60
.0094	.0082	15.8	.05190	.2278	Øį.	12.80
.0099	.cors	13.4	.00895	.0946	49	5.31
.0106	.0039	15.8	.02320	.1523	82	8.57
.0109	.0037	15.2	.०३५५०	.1562	29	2.78
.0110	.0039	16.2	.02410	.1552	38	3.74
.0110	.0139	20.2	.06690	.2625	83	14.75
.0131	.0213	21.6	.09870	.3242	10	17.60
.0123	.0366	22.7	.16100	.4012	W	22.50
.0132	.0323	21.7	.14900	.3860	N	21.60
.0135	.0226	14.7	.01.770	.1330	91	7.47
.0139	.0017	15.0	.01130	.1063	17	5.27
Fettke's 19	34 Jats (E 1s	Calculated)				
.0092	.0004	10.8	.00370	.0508	5.619	3.42
.0092	.0010	12.5	.00800	.0594	99	5.02
.0093	.0024	14.5	.01560	.1238		7.23
.0074	.0038	15.8	.02400	.1549	131	3.74
.0099	.0015	13.4	.01200	.1095	99	6.15
.0106	.0045	16.8	.02680	.1637	85	9.20
.0109	.0031	15.2	·05070	.1423	88	8.02
.0110	.0044	16.2	.02720	.1649	10	9.26
.6110	.0154	50.5	.07630	.2762	99	15.50
.0121	.0225	21.6	.10400	.3225		18.10
.0128	.0293	22.7	.13130	*3624	99	20.40
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.0135	.0026	14.7	.01770	.1330	13	7.46
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.0139	.0029	15.0	.01930	.1339	10	7.80
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Pettke's 19	38 Data (X is	Experimental)				
Pettke's 19	38 Sata (% is	Experimental)	.02090	.1446	5.619	9.12
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.0041 .0042 .0042 .0042	.00278 .00047 .00181 .00550	13.3 12.1 12.8 14.3	.02090 .00389 .01110 .03840	.1446 .0624 .1187 .1960	5.619	9.12 3.50 6.63 11.00
.0041 .0042 .0042 .0042 .0042	.00278 .00047 .00181 .00550 .00912	13.3 12.1 12.8 14.3 15.2	.02090 .00389 .01410 .03840 .05930	.1446 .0624 .1187 .1960 .2445	5.619	0.12 3.50 6.63 11.00 13.72
.0041 .0042 .0042 .0042	.00278 .00047 .00181 .00550	13.3 12.1 12.8 14.3	.02090 .00389 .01110 .03840	.1446 .0624 .1187 .1960	5.619	9.12 3.50 6.63 11.00

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TABLE II Fettke's 1934 Sata (K is Experimental)

(T)	(5)	(3)	(1)	(5)	(6)	(7)
rain Sise	Paraeability	Porosity	1 - #	(\$/x)\$	pacific	1/3
d (om)	K (Darcies)	/ (per cent)			(carl)	
92	.0005	10.8	.392	14.70	3000	108.
.0092	.0005	12.5	.875	15.61	10,170	108.
.0093	.0034	14.5	.855	6.53	4980	107.
.0094	.0082	15.8	.842	4.39	3700	106.
.0079	.0012	13.4	.865	10.56	7350	101.
.0106	.0039	16.8	.832	6.57	5970	94.
.0109	.0037	15.2	Bull.	6.41	5160	91.
.0110	.0039	26.2	.338	0.45	3610	91.
.0110	.0139	20.2	.793	3.81	4340	91.
.0121	.0213	51.6	.78u	3.18	3950	82.
.0128	.0366	22.7	.773	2.49	3290	73.
.0132	.0323	21.7	.783	2.59	3230	75.
.0135	.0026	14.7	.853	7.52	5825	74.
.0139	.0017	15.0	.850	9.39	7460	71.
ettke's 19	34 Peta (Kis	Calculated)	onder of the second of the sec			is the construction of the Transport
.0072	.0004	10.8	.372	16.43	0.960	108.
	500 E 1 2 1 3 1 3	12.5	.875	11.18	7175	103.
	.0010			attitute aftitibilities.	COLUMN CO	50 AL SE
.0093	.0024	14.5	.355	7.77	5930	
.0093	.002h .0038	14.5 15.8	.855	0.45	5450	106.
.0093 .0094 .0099	.0024 .0038 .0015	14.5 15.8 13.4	.355 .342 .866	0.45 9.46	51,50 6580	106,
.0093 .0094 .0099 .0106	.002k .0038 .0015 .0045	14.5 15.8 13.4 16.8	.855 .842 .866 .832	6.45 9.46 6.12	5450 6580 5630	106,
.0093 .0094 .0099 .0166 .0109	.002k .0038 .0015 .0045 .0031	14.5 15.8 13.4 16.8 15.2	.855 .842 .866 .832	6.45 9.46 6.12 7.00	5450 6580 5630 5650	106, 101, 94, 91,
.0093 .0094 .0099 .0166 .0109	.002k .0038 .0015 .00k5 .0031 .00kk	14.5 15.8 13.4 16.8 15.2 16.2	.855 .842 .856 .832 .848 .838	6.45 9.46 6.12 7.00 6.07	5450 6580 5630 5650 5275	106, 101, 94, 91,
.0093 .0094 .0099 .0106 .0109 .0110	.002k .0038 .0015 .0045 .0031 .00kk	14.5 15.8 13.4 16.8 15.2 16.2 20.2	.855 .842 .866 .832 .848 .838	6.45 9.46 6.12 7.00 6.07 3.62	5450 6580 5630 5650 5275 4120	106. 101. 94. 91. 91.
.0093 .0094 .0099 .0166 .0109 .0110 .0110	.002k .0038 .0015 .0045 .0031 .004k .015k	14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6	.855 .842 .866 .832 .848 .838 .798 .784	5.45 9.46 6.12 7.00 6.07 3.62 3.10	5450 6580 5630 5650 5275 4120 3840	106. 101. 94. 91. 91. 91.
.0093 .0094 .0099 .0166 .0109 .0110 .0121	.002k .0038 .0015 .0045 .0031 .0044 .015k .0225	14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6 22.7	.855 .842 .866 .832 .848 .838 .798 .784	5.45 9.16 6.12 7.00 6.07 3.62 3.10 2.76	5450 6580 5630 5650 5275 4120 3840 3645	106. 101. 94. 91. 91. 91. 62. 78.
.0093 .0094 .0099 .0166 .0109 .0110 .0121 .0128	.002k .0033 .0015 .0045 .0031 .0044 .015k .0225 .0298	14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6 22.7 21.7	.855 .842 .866 .832 .848 .838 .798 .784 .773	5.45 9.46 6.12 7.00 6.87 3.62 3.10 2.76 3.07	5450 6580 5630 5650 5275 4120 3840 3645 3830	106. 101. 94. 91. 91. 92. 78. 75.
.0092 .0094 .0099 .0166 .0109 .0110	.002k .0038 .0015 .0045 .0031 .00kk	14.5 15.8 13.4 16.8 15.2 16.2	.855 .842 .866 .832 .848 .838	6.45 9.46 6.12 7.00 6.07 3.62	5450 6580 5630 5650 5275 4120	
.0093 .0094 .0099 .0106 .0109 .0110 .0110 .0121 .0128 .0135 .0135	.002k .0033 .0015 .0045 .0031 .004k .015k .0225 .0296 .0331 .0026 .0029	14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6 22.7 21.7 14.7 15.0	.855 .842 .866 .832 .848 .838 .798 .784	5.45 9.46 6.12 7.00 6.07 3.62 3.10 2.76 3.67 7.52	5450 6580 5630 5650 5275 4120 3840 3645	106 101 91 92 92 92 91 76 75 71
.0093 .0094 .0099 .0166 .0109 .0110 .0110 .0121 .0128 .0132 .0135 .0139	.002k .0033 .0015 .0045 .0031 .0025 .025 .0298 .0331 .0026 .0029	14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6 22.7 21.7 14.7 15.0	.855 .842 .866 .832 .848 .838 .798 .784 .773 .783 .853	5.45 9.46 6.12 7.00 6.07 3.62 3.10 2.76 3.07 7.52 7.19	54.50 6580 5630 5650 5275 4120 3840 3645 3830 5025 5700	106 101 94 91 91 91 102 78 75 74 71
.0093 .0094 .0099 .0166 .0109 .0110 .0121 .0128 .0132 .0135 .6139	.002k .0033 .0015 .0015 .0031 .0025 .0225 .0296 .0331 .0026 .0029	14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6 22.7 21.7 14.7 15.0	.855 .842 .866 .832 .848 .838 .798 .784 .773 .783 .853 .850	5.45 9.46 6.12 7.00 6.07 3.62 3.10 2.76 3.07 7.52 7.19	5450 6580 5630 5650 5275 4120 3840 3645 3830 5025 5700	106. 101. 94. 91. 91. 92. 78. 75. 74. 71.
.0093 .0094 .0099 .0106 .0109 .0110 .0121 .0128 .0132 .0135 .0139	.002k .0033 .0015 .0045 .0045 .0031 .015k .0225 .0298 .0201 .0029	14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6 22.7 21.7 14.7 15.0	.855 .842 .866 .832 .848 .838 .798 .784 .773 .783 .853 .850	6.45 9.h6 6.12 7.00 6.07 3.62 3.10 2.76 3.07 7.52 7.19	5450 6580 5630 5650 5275 4120 3840 3645 3830 5025 5700	106, 101, 94, 91, 91, 92, 78, 75, 74, 71,
.0093 .0094 .0099 .0166 .0109 .0110 .0121 .0128 .0132 .0135 .0139	.002k .0033 .0015 .0015 .0015 .0015 .0015 .0015 .0015 .0025 .0298 .0298 .0298 .0200 .0020	14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6 22.7 21.7 14.7 15.0 Experimental)	.855 .842 .866 .832 .848 .838 .798 .794 .773 .783 .853 .850	6.91 16.06 6.45 9.h6 6.12 7.00 6.07 3.62 3.10 2.76 3.07 7.52 7.19	5450 6580 5630 5650 5275 4120 3840 3645 3930 5025 5700	106. 101. 94. 91. 92. 78. 75. 74. 71.
.0093 .0094 .0099 .0166 .0109 .0110 .0121 .0128 .0132 .0135 .0139	.002k .0033 .0015 .0045 .0031 .004k .015k .0225 .0298 .0031 .0026 .0029	14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6 22.7 21.7 14.7 15.0 Experimental)	.855 .842 .866 .832 .848 .838 .798 .784 .773 .783 .853 .850	6.45 9.46 6.12 7.00 6.07 3.62 3.10 2.76 3.07 7.52 7.19	5450 6580 5630 5650 5275 4120 3645 330 5325 5700	106. 101. 94. 91. 91. 92. 78. 75. 74. 71.
.0093 .0094 .0099 .0166 .0109 .0110 .0110 .0121 .0128 .0135 .0135 .0139	.002k .0033 .0015 .0015 .0015 .0031 .0025 .0298 .0231 .0026 .0029	14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6 22.7 21.7 14.7 15.0 Experimental)	.855 .842 .866 .832 .848 .838 .798 .784 .773 .783 .853 .850	6.45 9.46 6.12 7.00 6.07 3.62 3.10 2.76 3.67 7.52 7.19	5450 6580 5630 5650 5275 4120 3840 3645 3830 5025 5700	106. 101. 94. 91. 91. 91. 92. 78. 75. 74. 71.
.0093 .0094 .0099 .0166 .0109 .0110 .0110 .0121 .0128 .0135 .0135 .0139	.002k .0033 .0015 .0015 .0015 .0031 .0025 .0298 .0231 .0026 .0029 .00278 .00278 .00278 .00131 .00550 .00912 .00534	14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6 22.7 21.7 14.7 15.0 Experimental)	.855 .842 .866 .832 .848 .838 .798 .784 .773 .783 .853 .850	6.45 9.46 6.12 7.00 6.07 3.62 3.10 2.76 3.07 7.52 7.19 6.91 16.06 8.41 5.10 4.08 5.18	5450 6580 5630 5650 5275 4120 3840 3645 3930 5025 5700 	244 238 238 227
.0093 .0094 .0099 .0106 .0109 .0110 .0121 .0128 .0135 .0135 .0139	.002k .0033 .0015 .0015 .0015 .0031 .0025 .0298 .0231 .0026 .0029	14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6 22.7 21.7 14.7 15.0 Experimental)	.855 .842 .866 .832 .848 .838 .798 .784 .773 .783 .853 .850	6.45 9.46 6.12 7.00 6.07 3.62 3.10 2.76 3.67 7.52 7.19	5450 6580 5630 5650 5275 4120 3840 3645 3830 5025 5700	107. 106. 101. 94. 91. 91. 92. 78. 75. 74. 71.

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TABLE IV

(K is experimental)

Permeability K (Services)	Grain Size d (cm)	(3) K/d ²	(4) K/d ² / 1.013 X 10 ⁸	(5) Perceity \$ (per cent)
.0005	.0092	5.91	5.83 × 10 ⁻⁸	10.8
.0005	.0092	5.91	5.03	12.5
.0034	.0093	39.30	38.70	14.5
.0032	.0094	92.80	91.60	15.8
.0012	.0099	12.25	12.08	13.4
.0039	-0106	34.80	34.30	16.8
.0037	.0109	31.20	30.75	15.2
.0039	.0110	32.20	31.75 113.50	15.2 20.2
.0213	.0121	146.00	1kh.00	21.6
.0366	.0123	223.00	220.00	22.7
.0323	.0132	185.60	183.00	21.7
.0026	.0135	14.30	14.10	14.7
.0017	.0139	8.80	8.68	15.0
LE 15 CELEVILATA	rd)			
.0004	.0092	4.72	11.66 x 10 ⁻³	10.8
*0010 *000ft	.0092	11.30	11.64	12.5
.0004	.0092		11.66 x 10 ⁻³ 11.64 27.40 42.40	
.0010 .000ft	.0092 .0092 .0093	11.80 27.80	11.64 27.40	12.5
.0004 .0010 .0024 .0038 .0015	.0092 .0092 .0093 .0094 .0099	11.80 27.80 43.00 15.30 40.20	11.64 27.40 42.40 15.10 39.60	12.5 14.5 15.8 13.4 16.8
.0004 .0010 .0024 .0038 .0015 .0045	.0092 .0092 .0093 .0094 .0099 .0106 .0109	11.30 27.30 43.00 15.30 40.20 26.10	11.64 27.40 42.40 15.10 39.60 25.70	12.5 14.5 15.8 13.4 16.8 15.2
.0004 .0010 .0024 .0038 .0015 .0045 .0031	.0092 .0092 .0093 .0094 .0099 .0106 .0109	11.30 27.30 43.00 15.30 40.20 26.10 36.40	11.64 27.40 42.40 15.10 39.60 25.70 35.90	12.5 14.5 15.8 13.4 16.8 15.2 16.2
.0004 .0010 .0024 .0038 .0015 .0045 .0031 .0044	.0092 .0092 .0093 .0094 .0099 .0106 .0109 .0110	11.30 27.30 43.00 15.30 40.20 26.10 36.40 128.20	11.64 27.40 42.40 15.10 39.60 25.70 35.90 126.60	12.5 14.5 15.8 13.4 16.8 15.2 16.2 20.2
.0004 .0010 .0024 .0038 .0015 .0045 .0031 .0044	.0092 .0092 .0093 .0094 .0099 .0106 .0110 .0110	11.30 27.80 43.00 15.30 40.20 26.10 36.40 128.20 154.00	11.64 27.40 42.40 15.10 39.60 25.70 35.90 126.60 152.00	12.5 14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6
.0004 .0010 .0024 .0038 .0015 .0045 .0031 .0044 .0154	.0092 .0092 .0093 .0094 .0099 .0106 .0109 .0110 .0121	11.30 27.80 43.00 15.30 40.20 26.10 36.40 128.20 154.00 132.00	11.64 27.40 42.40 15.10 39.60 25.70 35.90 126.60 152.00 179.50	12.5 14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6 22.7
.0004 .0010 .0024 .0038 .0015 .0045 .0031 .0044	.0092 .0092 .0093 .0094 .0099 .0106 .0110 .0110	11.30 27.80 43.00 15.30 40.20 26.10 36.40 128.20 154.00	11.64 27.40 42.40 15.10 39.60 25.70 35.90 126.60 152.00	12.5 14.5 15.8 13.4 16.8 15.2 16.2 20.2 21.6

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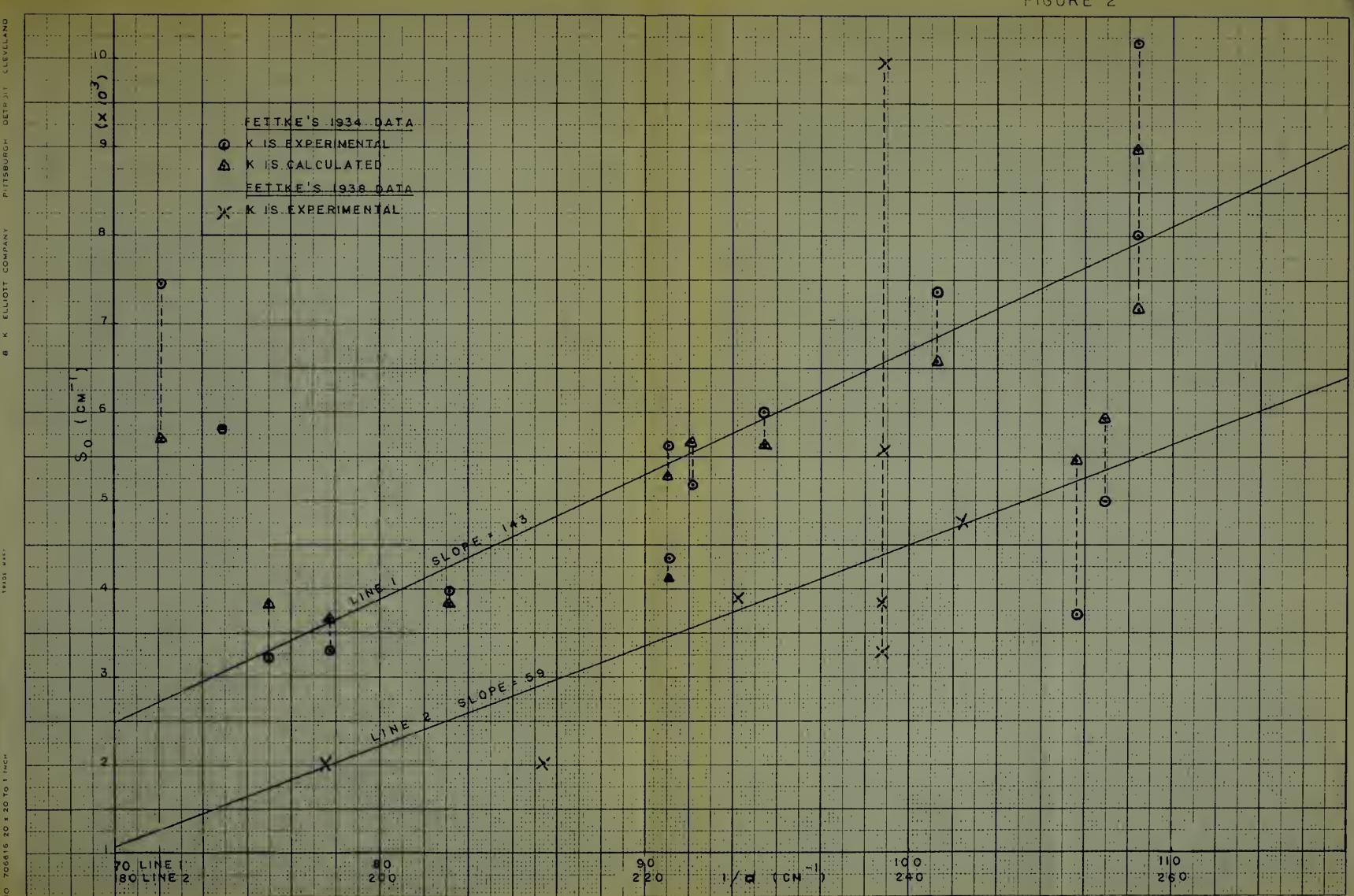
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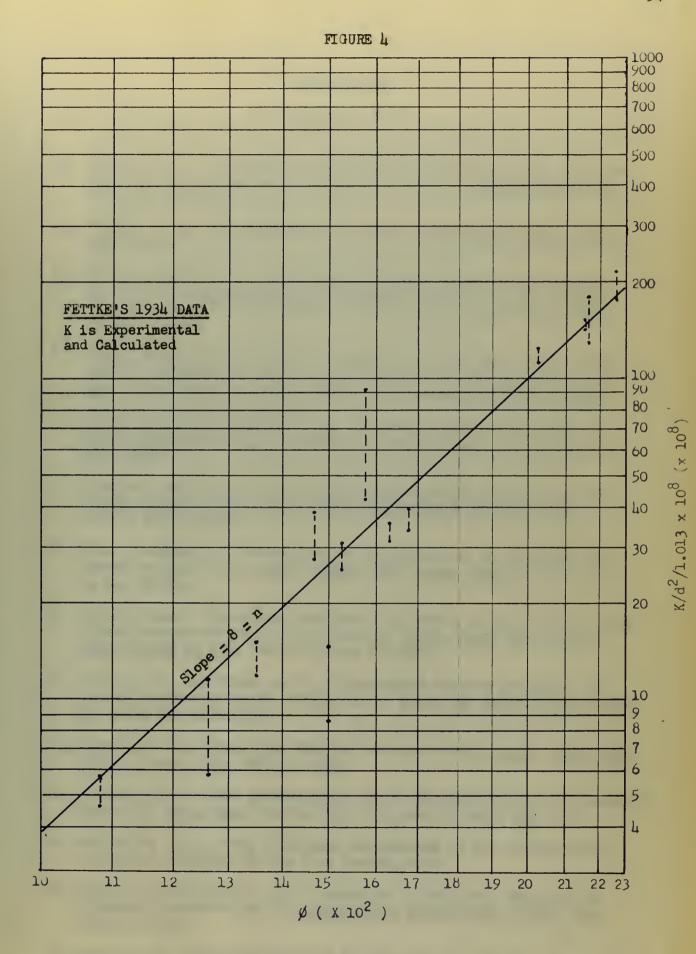
APPENDIX II













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